Morphological and Microstructural Characterization of Organoclays from Low Smectite Containing Clays Materials

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Morphological and Microstructural Characterization of Organoclays from Low Smectite Containing Clays Materials

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Authors’ contributions

This work was carried out in collaboration between all authors. All authors read and approved the final manuscript.

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ABSTRACT

Two low smectite-containing clay materials were modified, using hexadecyltrimethylammonium bromide (HDTMAB) as intercalating agents, under very mild experimental conditions, to investigate their potentials as suitable organoclays for industrial and environmental applications. Changes in the general morphological and microstructural characteristics were studied by X-Ray Diffraction (XRD) analysis, Thermogravimetric Analysis (TGA), Brauner Emmet Taylor (BET) Analysis, Fourier Transform Infrared (FTIR) Spectroscopy and Scanning Electron Microscopy (SEM) before and after modification. The shift in XRD reflections after intercalation indicated that the HDTMA chains adopt monolayer and bilayer arrangements within the clay interlayers and were largely dependent on the reaction time and surfactant loading. This resulted in decreased specific surface area and increased pore sizes in the organoclay samples. The presence of typical CH stretching bands in
the new materials was confirmed by FTIR analysis. Mass loss from TGA showed that the new materials were thermally stable, the amount of organic modifier in the organoclays were in good agreement with the theoretically calculated contents. The intercalated materials were hydrophobic, stable, biocompatible and were suitable materials for several industrial and environmental applications.

Keywords: Clays; HDTMAB; interlayers; organoclays; intercalation; smectites.

1. INTRODUCTION

The need to develop and study clay materials with notable properties and of useful applications such as environmental remediation [1–4], water purification [1,5,6], rheological additives [7–10], catalysis [11], food and agriculture [12–16], drugs [10,16–18] and nanohybrid precursors [19–22] has been on the rise over the past decade. Awareness and use of various types of clays and their beneficial effects can be traced down to the ancient Mesopotamians, Chinese, Indian and Egyptian civilizations to as far back as three to five hundred years ago [23]. Clays find a wide range of applications, in various areas of science due to their natural abundance and the propensity with which they can be chemically and physically modified to suit practical technological needs [24]. Clay is composed mainly of silica, alumina and water, frequently with appreciable quantities of iron, alkalis and alkali earths [25]. Clay minerals are natural environment-friendly materials with high specific surface areas and now widely used for the adsorption and removal of the organic pollutants [3,5,26–29]. They have been well known to mankind from the earliest days of civilization, with variability in their constitution depending on their groups and sources [18,25]. Clay minerals are essentially hydrous aluminosilicates with very fine particle sizes. Due to an isomorphic substitution within the layers, the clay mineral layer is negatively charged, which is counterbalanced by the exchangeable cations such as Na⁺, Ca²⁺ in the interlayers [30]. The hydration of the inorganic exchange ions present in clays and the nature of Si-O groups impart a hydrophilic nature to the mineral surfaces and this property makes them easy to absorb water but very difficult to disperse in organic matrices and hence renders the clay ineffective for interaction with aliphatic and relatively hydrophobic compounds [31,32]. The most studied methods for solving this problem generally involve the replacement of these exchangeable cations with an organic group. The process usually involves cation exchange with the surface and interlayer clay cations or a solid state reaction which usually involves an ion-dipole interaction. The organic group is typically intercalated between the clay interlayer spaces and (or) adsorbed on the surface where the inorganic ions were previously attached. These types of modified clays are most often referred to as organoclays. In organoclay synthesis, the smectites clays are widely used, by virtue of their high cation exchange capacity, excellent ability to swell, high platelet aspect ratio and ease with which their surface can be modified [32,33]. Ion exchange with alkylation ammonia ions is the preferential method to prepare organoclays [9] and was employed in this study.

In this study, two natural clays materials with low smectite contents were selected, modified and investigated for their potential as suitable starting materials for organoclays. Low smectite-containing clays have an original low smectite content (ranging from near zero to approximately 15% smectite content). This work was aimed at modifying the materials by selective intercalation of their smectite compositions, under mild experimental conditions. How these modifications influenced some of the important properties of the materials were investigated. The synthesis of organoclays is dependent on the mechanisms of the reactions that the clay minerals can have with the organic compounds used. The effect of time and surfactant loading were specifically studied to give insight on the mechanism behind the transformation and also to address how the materials changed in their properties and morphology when exposed to hexadecyltrimethylammonium (HDTMA) bromide, a popular intercalant in clay science.

2. MATERIALS AND METHODS

Two natural low smectite-containing clay samples sourced within Nigeria [Adamawa (9°33'N 12°43'E) and Yobe (12°00'N 11°30'E) States], coded Sample A and Sample C respectively, were collected from the National Geosciences Research Laboratory, Kaduna, Nigeria. All samples were pulverized, sieved (<200 µm), and stored in dry plastic bags prior to modification. HDTMA bromide (Sigma-Aldrich)
was used to carry out the cation exchange reaction in the synthesis of the organoclays using the simple hydrothermal method [30,31], under very mild reaction conditions.

Prior to hydrothermal modification, the clay samples were purified using the synthetic methods described in Pedro et al., [34], where the clay samples were first dispersed in distilled water to separate unwanted fractions by sedimentation and filtration based on Stokes’ law of settling in suspension. The collected samples, like with most raw clays contained a mixture of several mineral phases as impurities. Smectites, like most clay minerals, by their very nature, are extremely fine grained, and it is this attribute that is utilized to separate them from more coarse-grained minerals with which they commonly coexist (quartz, feldspar, calcite and other impurities). The dispersed clay samples were collected and treated with Dithionite-Citrate-Bicarbonate (DCB) to remove all iron oxide and hydroxide in the clay suspension. The DCB solution was removed by centrifugation and the clay samples washed severally until pH was 7. The treated clay samples were finally treated with hydrogen peroxide and sodium acetate solution to remove organic substances from the clay as explained in Manocha et al. [31]. The treated samples were dried at 105°C and labelled as Sample 1 and Sample 3 for Sample A and Sample C respectively.

For Intercalation experiments, the treated samples (Sample 1 and Sample 3) were placed in separate 200 ml beakers and heated in a water bath until they reached the desired temperature. To the samples in the beakers, selected concentrations based on the CEC of the salt sample were added. The mixture was stirred until it reached the desired time for the hydrothermal process to complete. The samples were then removed from the beaker and centrifuged and washed with hot distilled water until all loose Br atoms were removed from the organoclay solution. This was confirmed using AgNO₃ test, as no precipitates were formed in the collected supernatant solution. The centrifuged organoclay was dried in the oven at 85°C, pulverized, sieved below 200 µm and then kept in a desiccator for subsequent experimentation. The procedure was repeated by varying the time (2.5, 5, 7.5, 10 and 12.5 hours), on Sample A and by varying the surfactant concentrations (5, 10, 20, 30, 40 and 50 percent w/w material) on Sample C.

The Cation Exchange Capacity (CEC) of the raw and purified clay materials were determined using the barium chloride compulsive exchange method [35]. The surface area, porosity and pore sizes of the clays and their corresponding HDTMA organoclays were measured from N₂ sorption isotherm at 77 K using a Micrometrics Tristar II BET analyser. The samples were degassed under flowing nitrogen at 250°C for 4 hours using the Smart Prep™ 065. The degassing unit removes adsorbed contaminants from the surface and pores of a sample in preparation for Brauner Emmet Taylor (BET) analysis. The chemical compositions of the clay minerals before and after modification were measured by X-ray fluorescence spectroscopy (XRF) using an Oxford Instrument (X-supreme 8000) with a tungsten X-ray tube and powder sample holder. Concentrations were expressed in percentages (%).

X-Ray Diffraction (XRD) analysis gives information about the interlayer spacing of the materials which is very important for explaining the intercalation and configuration of organic surfactant into the material and also enables the determination of the expansion as well the crystal size of the clay [36]. X-ray Diffractometer (XPERT-PRO, PW3064, Copper Kα radiation at 40 kV/40 mA, with a goniometer velocity of 2°/min and a step of 0.05°, with 2θ ranging from 0° to 89°) was used to probe the crystal lattice structure and interlayer space of the clay and organoclay powders. To determine whether a chemical reaction took place between organic compounds and the clay samples and also to investigate the position, manner of arrangement and presence of characteristic organic groups in the sample’s structure, Fourier Transform Infrared (FTIR) analysis was carried out. FTIR characterization has been proven to be an effective method for exploring the microstructures of organoclays [37–40]. The FTIR spectra were recorded on a Shimadzu FTIR-8400s Spectrophotometer using KBr, with 20 scans collected for each measurement over the spectral range of 4000-400 cm⁻¹ with a resolution of 4 cm⁻¹. Each sample (2.0 mg) was mixed thoroughly with 150 mg of spectroscopic grade KBr and dried in an oven at 120°C for about 12 hours before analyses. All spectra were presented without baseline correction or normalization. Since FTIR spectroscopy measures the molecular vibrations and determines the molecular structure of the organoclays, any change in the structure of the
organoclay will be observed through alterations in the infrared spectra [17,36].

To determine the surface morphology before and after modification, Scanning Electron Microscopy (SEM) was carried out using a Phenom™ X-pro Scanner, with 5 KV of accelerating voltage. The samples were sputter coated with thin conductive gold before analysis. Thermogravimetric Analysis (TGA) and Differential Thermal Analysis (DTA) were used to investigate the thermal and energy profile of the clay and organoclays materials. Thermal studies play an essential role in determining both technological applications and processing conditions of the clay materials and can lead to new insights into the structure of the intercalated clay [37,41]. A Mettler-Toledo TGA/DSC1 Star® (Mettler-2012) instrument with a heating rate of 20°C/min in Nitrogen (N₂) atmosphere was employed in determining the thermal characteristics of the raw and modified clay samples. All of the samples were heated up to 500°C.

3. RESULTS

3.1 Analysis of Surface Properties and Morphology

N₂ adsorption data from BET analysis used to investigate the surface properties of the materials are presented in Table 1. Decreased values in the pore volumes and SₐBET of the samples were observed after modification. The observed reduction in the pore volume of Sample A and Sample C from 0.054 and 0.053 cm³/g respectively to 0.017 cm³/g after modification are due to the organic molecules filled in the clay interlayers. Carmody et al. [42] observed that drying materials may result in a significant and irreversible reduction in pore volumes and hence resulting to the internal SA of the materials being inaccessible to N₂ molecules. The measured specific surface areas of clay minerals are the sum of the surface areas of the interspaces (micropores), the intraparticle spaces (micro/mesopores inside the assembly of the silicate layers), and the inter-aggregate spaces (meso/micropores inside the assembly of particles [43,44]. The SₐBET of Sample A and Sample C which were 18.256 m²/g and 47.728 m²/g were also reduced to 5.113 and 5.323 m²/g in Samples 2 and 5 respectively. The decreased surface area of the organoclay samples, with respect to the values of their corresponding raw clay samples, is a consequence of the disappearance of micropores, which is caused by the blocking of the N₂ passage by the large HDMA cations in the clay interlayer, leaving only a small fraction of the interlamellar pores accessible [44,45]. The decrease in SₐBET of the organoclays largely depends on the packing density which may be predominantly occupied by long-chain HDMA⁺ in the interlamellar layer resulting in more serious pore blocking that inhibits the passage and adsorption of N₂ in the organoclay [42,46,47,48], since the HDTMA cations were too strongly bound to be physically replaced by N₂ molecules. These findings suggest that the HDTMA cations are introduced inside the interlayer of the clay mineral not only by cation exchange at the planar sites but also through the interaction with the aluminosilicate sheets [48,49]. The surface areas of the organoclays and raw clay samples are interrelated at the level of stacking of elementary layers in that the adsorption by reaction with the planar sites and the consequent formation of outer-sphere complexes is likely to occur [50]. Higher surface area of raw clay samples in comparison to that of their respective organoclay samples suggests that the latter samples may have a considerable amount of the trapped pores in the range of small and medium mesopores (20–400 Å). The amount of the trapped pores in the raw samples are lower than those of their corresponding organoclays which can be explained by the development of larger mesopores as a result of the arrangement of HDMA cations in the interlayer spaces of these samples. IUPAC recommendation for pores are classified as <2nm for micropores, 2nm<size<50nm for mesopores and >50nm for macropores. The average pore sizes of the samples showed that they were mainly mesoporous and had increased from 124.392 Å (Sample A) and 58.355 Å (Sample C) in the raw samples to 137.396 Å (Sample 2) and 148.449 Å (Sample 5) respectively, along with the observable decrease in pore volumes. Increased pore sizes suggested that the surfactants were intercalated into the interlayer spaces of the clays rather than being adsorbed/precipitated to/in interparticles and aggregate species. The insertion of HDMA⁺ cations, by the exchange process, into the raw clays, increased the number of medium and large mesopores and caused the significant decrease in the surface areas and pore volumes of the organoclays compared to the raw samples. The results indicated that the exchanged cations affect the surface characteristics of the clays in some manner that appears to be related to the size and interlayer arrangement of the exchanged ion in
the material samples [48]. These results are similar to the findings of other researchers [44] indicating that the materials can sufficiently be applied as functional organoclays.

### 3.2 Analysis of Crystal Structures by XRD

The characteristic d001 smectite reflections of the raw and the modified clay materials studied by XRD were summarized in Table 1. Reflections corresponding to the 001 reflections of smectite minerals at $^\circ 2\theta$ positions of approximately 7 $^\circ 2\theta$ or 12.5 Å were clearly indicated. Reflections relating to the presence of other mineral phases were an indication that the purification procedure was not absolute for the samples employed. In the diffractograms of Samples A and B, the characteristic reflections were at 7.6 $^\circ 2\theta$ and 5.8 $^\circ 2\theta$, which also correspond to interlayer spaces of 11.5 Å and 15.3 Å respectively. The effects of reaction time and surfactant dosage were also studied and are illustrated in Fig. 1 and Fig. 2.

<table>
<thead>
<tr>
<th>Properties</th>
<th>Raw clay samples</th>
<th>Purified clay samples</th>
<th>Organoclay samples</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>Sample A</td>
<td>Sample C</td>
<td>Sample 1</td>
</tr>
<tr>
<td>Modified from</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>$\text{Al}_2\text{O}_3$ (%)</td>
<td>29.141</td>
<td>25.797</td>
<td>32.324</td>
</tr>
<tr>
<td>$\text{SiO}_2$ (%)</td>
<td>65.538</td>
<td>53.002</td>
<td>60.8</td>
</tr>
<tr>
<td>$\text{TiO}_2$ (%)</td>
<td>2.211</td>
<td>1.52</td>
<td>1.29</td>
</tr>
<tr>
<td>$\text{Na}_2\text{O}$ (%)</td>
<td>0.077</td>
<td>2.959</td>
<td>1.26</td>
</tr>
<tr>
<td>$\text{MgO}$ (%)</td>
<td>0.523</td>
<td>1.831</td>
<td>1.043</td>
</tr>
<tr>
<td>$\text{Fe}_2\text{O}_3$ (%)</td>
<td>1.318</td>
<td>12.124</td>
<td>1.112</td>
</tr>
<tr>
<td>MC (%)</td>
<td>0.40</td>
<td>3.27</td>
<td>0.64</td>
</tr>
<tr>
<td>OMC (%)</td>
<td>0.54</td>
<td>4.00</td>
<td>0.00</td>
</tr>
<tr>
<td>Pore Volume (cm$^3$/g)</td>
<td>124.392</td>
<td>58.355</td>
<td>103.121</td>
</tr>
<tr>
<td>Pore Size (Å)</td>
<td>18.26</td>
<td>47.73</td>
<td>22.23</td>
</tr>
<tr>
<td>$\text{SA}_{\text{BET}}$ (m$^2$/g)</td>
<td>98.24</td>
<td>92.6</td>
<td>102.32</td>
</tr>
<tr>
<td>CEC (Meq/g)</td>
<td>11.6</td>
<td>15.3</td>
<td>11.6</td>
</tr>
</tbody>
</table>

MC stands for moisture content; OMC, Organic matter content; $\text{SA}_{\text{BET}}$, Bruner Emmet Teller Specific Surface Area; CEC, cation exchange capacity; and OC stands for the Organoclay samples tested.

Fig. 1. The effect of time on intercalation: XRD diffractograms of samples A, 2, 10, 11 and 13
After intercalation, the reflections shifted leftward from this first characteristic reflection towards lower angles, indicating an increase in the interlayer space of the materials. The clay sheets were bound with in-plane covalent bonds and therefore their crystal structures were stable. The change in the crystal structure was indicative of the expansion of the interlayers by the alkylammonium ion. The interlayer space of alkylammonium clays changes in a characteristic way with the length of the alkylammonium [9]. Clay mineral layers are held loosely with Van der Waals bonds and hence expansion of interlayers is commonly seen when water or organic molecules are introduced between the layers [39]. The layers of the clays were propped open upon swelling in water in the presence of the alkylammonium surfactant. After the ion exchange process, the cationic head groups of the intercalation agent molecule would preferentially reside at the layer surface and the aliphatic tail will radiate away from the surface expanding the d-value in comparison with that of the original clay sample [36,37]. This is a characteristic property of smectites.
Although exchange reaction were quite fast (about 40–120 min), Nguyen et al. [51] suggested that in order for cations to be stable between clay mineral layers, reaction solution should be kept in longer time. The modification process, carried out at 80°C using 50% wt. HDTMA was repeated for 2.5, 5, 7.5, 10 and 12.5 hours. An evident increase in interlayer space with increasing time from 11.5 Å for Sample A to 20.3, 16.5, 17.5, 18.1, and 19.6 Å for the modified Samples 2, 10, 11, 12 and 13 respectively were shown in Fig. 1. The modification process induced noticeable effects on interlayer space after the first 2.5 hours (16.52 Å), and the maximum interlayer space (20.3 Å) was obtained after 10 hours. No further increase in spacing was observed after the 10-hour process. Instead, there was a decrease in the interlayer space to 19.6 Å when the process was allowed to stay for 12.5 hours. The shift of the XRD reflections as time progressed can be attributed to re-orientation of the long organic salt chain between MMT layers into more condensed configurations [52,53].

3.3 Analysis of Microstructural and Chemical Properties by FTIR Analysis

The FTIR spectra of the clay samples and their resultant organoclays, within the range of 4000 to 400 cm⁻¹, are shown in Fig. 2 and Fig. 3. The infrared spectra of the unmodified clays, Sample A and Sample C, showed major bands that are characteristic of smectite clay minerals. Bands in the 600-400 cm⁻¹ region were attributed to Si-O and Al-O bending bands with interlayer deformation around 1630 cm⁻¹. The tetrahedral bending modes were near 430 cm⁻¹ for Si-O-Si and 540 for Si-O-Al [58].Si-O bending modes was observed at 790 and 690 cm⁻¹ indicating the presence of quartz in the materials [25]. The intense broad bands observed between 1100 and 1020 cm⁻¹ were those for also Si-O stretching bands of the smectites. The band in the 3600-3440 cm⁻¹ region were ascribed to OH stretching of hydroxyl groups present in the interlayer of the clay material by covalent attachment to the silane molecules. The spectra showed prominent bands between 3622 and 3695 cm⁻¹ indicating the possibility of hydroxyl linkage due to OH stretching band of structural OH group from Al-OH and Si-OH-Al stretching (interlayer) with the bending mode at 914 cm⁻¹ (and 922 cm⁻¹), which are the characteristics of dioctahedral smectites.

The FTIR spectra of the organoclays showing the effect of time and ammonium salt concentration are also observed in Fig. 2 and Fig. 3 respectively. In the spectra of the organoclays, bands in the 3700-3300 cm⁻¹ region were attributed to the O-H absorptions of the surfactant and those in the 3000-2800 cm⁻¹ were those for also Si-O vibration. The FTIR spectra of the organoclays showed prominent bands between 2922 and 2852 cm⁻¹ indicating the possibility of hydroxyl linkage due to OH stretching band of structural OH group from Al-OH and Si-OH-Al stretching (interlayer) with the bending mode at 914 cm⁻¹ (and 922 cm⁻¹), which are the characteristics of dioctahedral smectites.
The IR bands that appeared in the 3000-2800 cm\(^{-1}\) region were the signatures of CH\(_2\) infrared absorption bands. These bands are indicative of the ordering (gauche/trans conformer ratio), packing density of the surfactants in organoclays and the interactions between the alkyl chains [40]. The bands at 2922 cm\(^{-1}\) were also due to asymmetric C-H stretching band of CH\(_2\) aliphatic and the absorption bands at 2852 cm\(^{-1}\) corresponds to symmetrical CH\(_2\) stretching [59]. The presence of CH\(_3\) and CH\(_2\) transmission bands in organoclays confirmed intercalation while the systematic decrease or weakening of intensity of OH (free) stretching bands confirmed that the materials became more and more hydrophobic with increasing alkylammonium ion concentrations. According to previous reports [29,41,60–62], the CH\(_2\) infrared absorption bands on FTIR patterns, especially the \(\nu_{as}\) (CH\(_3\)) mode, are sensitive to the packing density of the adsorbed surfactants on clays, and the absorption frequencies of \(\nu_{as}\) (CH\(_3\)) are believed to be inversely related to the packing density of the adsorbed surfactants [40]. It is evident that there was a shift for these bands towards lower wavenumbers as the surfactant loading rate increased from 25% to 70% (w/w). The results supported the evidence from the XRD results, that increasing surfactant packing density causes additional expansion of the clay interlayer spaces. The results of this study agree with the findings of Zhu et al. [40] and are attributed to the progressively developing conformation of the adsorbed surfactant molecules from low packing density and ordering (liquid like) to high packing density and ordering (solid like). Bands due to asymmetric and symmetric modes did not appear in the case of unmodified clays due to the absence of organic molecules therein. The peaks at 1660 - 1620 cm\(^{-1}\) were due to H-O-H bending band and the intensity reflected the amount of water in the interlayer [52]. The IR absorption band around 1480 cm\(^{-1}\) is normally ascribed to the antisymmetric angular deformation or bending mode of the head ((CH\(_3\))\(_3\)N\(^+\)) methyl group of the HDTMA\(^+\) and has been noted to be sensitive to the extent of disorder and packing of the head group [28,52].

### 3.4 Analysis of Microstructural Properties by SEM

Fig. 5 to Fig. 8 show the SEM images of the clay samples before and after modification. The original samples (Samples A and C) had massive, curved plates and bulky flakes with aggregated morphology. On the other hand, the modified samples (Samples 2 and 5) were discrete and gathered together much easier than those of the raw samples. They were characterized by several small and scattered platelets that are relatively flat. The modified samples showed a more heterogeneous agglomeration in terms of dimension and particle size (polydispersity), comprising many irregular lamellae very likely because of the interactions
between the alkyl groups of surface exchanged molecules belonging to different particles. The scattered platelets imply an increase in the interplanar distance by the presence of the intercalated alkylammonium ions that tends to transform into local clusters with a high packing density. It is also of importance to reveal that although the organoclay materials showed a significant change in morphology compared to their original nature, there were not many morphologic differences observed despite the obvious variation observed in the XRD and FTIR measurements. The retention of some characteristic properties of the natural clay materials after giving the chemical treatment can be distinctly observed.

Fig. 4. FTIR spectra of samples C, 15, 16 and 17

Fig. 5. SEM micrograph of sample A
3.5 Thermoanalytical Characterization and Degradation Behaviour

The TGA and DTG profiles of the raw and modified samples are shown in Fig. 8 to Fig. 11. Fig. 8 and Fig. 9 display the results of the raw clay samples while Fig. 10 and Fig. 11 display those of the modified samples.

At a given temperature the clays’ and organoclays’ mass losses are directly related to the rate of the material decomposition process [37]. The way in which the mass loss steps occur provides information on the properties of the inserting molecules. This study explains a four step mass loss covering 50 to 500°C to enable a better understanding of the sample characteristics.
The first step forms the ambient to 110°C temperature range and is attributed to desorption of water from the clay [63]. Free water (water between particles and adsorbed on the external surfaces of crystals) are released at temperatures of below 110°C. This explains the hydrophobic character and at the same time the moisture contents of the materials. The thermogram recorded an initial sample mass of 2.52 mg for Sample A which did not show any major decrease in the first 110°C. Sample C decreased in mass by about 3.27% (0.28 mg).
from an initial sample mass of 8.46 mg. The DTG of Samples C shows a deep broad through with a maximum at 90°C. This indicates that the raw samples were not thermally stable within the first step. This was only possible due to the loss of physisorbed and chemisorbed water [64]. In the organoclay samples, the initial masses of Samples 2 and 5 were 7.98 mg 4.93 mg respectively. There were no obvious changes in the masses of Samples 2 and 5 from the thermograms.

Fig. 10. Thermogram of sample C

Fig. 11. Thermogram of sample 2
The second step occurred from 110°C to 200°C temperature range and is assigned to the dehydration of the hydrated cations in the interlayers of the samples [33,63]. At 200°C, the masses of Samples A, C, 2 and 5 were observed to have been reduced by about 0.5% (0.00 mg), 3.0% (0.25 mg), 0.5% (0.04 mg) and 1.0% (0.05 mg) respectively. The DTG scans in Sample A continued fluctuating as the temperature increased while that of Sample C showed a continual increase at 0.64 mg/min. The organoclay samples did not show any clear change in the DTG profile as the temperature increased in the second step. This indicated an improvement in the thermal stability and hydrophobicity of the modified materials. The proof of hydrophobicity is reflected in the thermal stability of the organoclay samples within the first 200°C on the DTG scans.

The third mass loss step occurred within the temperature range of 200°C and 350°C and explains desorption and decomposition of organic molecules from the surface of the samples. Samples A and C were reduced from their initial masses by a total of 0.3% (0.01 mg) and 3.8% (0.32 mg). The DTG profile of the Samples A and C showed relatively stable but fluctuating patterns in this step. In the organoclay samples, this step corresponded to the chemical decomposition of the bonded structure of the organic modifier and indicated the extent of intercalation as a percentage of the modifier present in the materials. This explains two types of bonding of surfactant molecules in the organoclays. One type of bonding is to the silica surface and the second to other surfactant molecules. Sample 2 was reduced by 5.5% (0.44 mg) and Sample 5 by 18.0% (0.89 mg). The DTG profiles of Fig. 10 and Fig. 11 show an inflection from stability at 200°C which was the onset of the decomposition of the organic matter with the maximum decomposition around 250°C – 275°C, which was close to the melting point (253°C) of HDTMAB.

The final mass loss step occurred within the temperature range of 350°C - 500°C and is attributed to the loss of structural hydroxyl groups from within the clay. The thermogram of the samples showed a continuous decrease 350°C to 500°C for all the samples. At 500°C the samples had reduced totally by 1.72%, 7.2%, 10.5% and 22.8% for the Samples A, C, 2 and 5 respectively.

4. CONCLUSION

Raw clay materials with low smectite contents and varying compositions of impurities have been effectively modified into more organophilic, thermally stable and functional organoclays.
materials by hydrothermal cation exchange with HDTMA\textsuperscript{+} for the purpose of nano-hybrid precursors and sorbents for environmental remediation. The HDTMA-exchanged clay materials showed not only morphological features such as decreased specific surface area and the irregular, wavy surface with curled edges but also structural features such as larger d-values and lattice distortion. The XRD results proved that an increase in the smectite interlayer space of the materials as a result of the intercalated HDTMA cations and that the HDTMA arrangement strongly depended on the duration of intercalation and on the packing density of surfactant within the clay galleries. This can facilitate the entry of the host organic molecules into the clay mineral interlayer space in further applications. The thermograms have revealed that these materials can be changed from being hydrophilic to hydrophobic, which is also a characteristic of successfully intercalated smectites. The intercalated structures thus formed have better heat stabilities. The selective modification of the smectites present in the clay mixture, therefore, influenced a change in the overall morphological and structural characteristics of the materials. These findings are important and vital to the preparation of low-cost and biodegradable organoclays for industrial and environmental applications.

**COMPETING INTERESTS**

Authors have declared that no competing interests exist.

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